



# Long-term effects of lime and phosphogypsum application on tropical no-till soybean–oat–sorghum rotation and soil chemical properties

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## ABSTRACT

Root growth, nutrition and crop yield can be affected by soil chemical modifications caused by superficial limestone and phosphogypsum application in a no-till system. Using this approach, this study was conducted in southeastern Brazil, continuing an experiment that has been on-going since 2002 with the objective of evaluating the residual effects of the surface application of lime and phosphogypsum on the soil chemical characteristics and the root growth, nutrition and yield of soybean, black oat and sorghum in a dry winter region cultivated in 2008/2009 and 2009/2010. The experimental design was a randomized block with 4 replications. The treatments were applied in November 2004 and were as follows: original conditions, limestone application ( $2000 \text{ kg ha}^{-1}$ ), phosphogypsum application ( $2100 \text{ kg ha}^{-1}$ ), and limestone ( $2000 \text{ kg ha}^{-1}$ ) + phosphogypsum ( $2100 \text{ kg ha}^{-1}$ ) application. Superficial liming with or without phosphogypsum reduced the surface and subsurface soil acidity 5 years after application in the no-till system. The movement of  $\text{Ca}^{2+}$  and  $\text{Mg}^{2+}$  from the surface layer into the subsoil over time was evident. The phosphogypsum application associated with liming increased the  $\text{Ca}^{2+}$  levels throughout the soil profile. Liming maintained high levels of  $\text{Mg}^{2+}$  throughout the soil profile with or without phosphogypsum application. The organic matter content increased with liming with or without phosphogypsum, indicating that in the long term, these practices can increase the C accumulation. Phosphogypsum application had a residual effect on the  $\text{SO}_4^{2-}$  levels, and high sulphate concentrations were observed in the subsoil after 5 years. Superficial liming improved crop nutrition and, when associated with phosphogypsum, increased Ca absorption by soybean and sorghum, as reflected in the increased yields of these crops.

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## 1. Introduction

Soil acidity is one of the factors limiting crop production, especially in tropical areas such as Brazil. The total area that is affected by acidity is estimated to be 4 billion ha, representing ~30% of the cropland worldwide (Sumner and Noble, 2003). The affected areas are mostly within countries located in the tropics; they comprise primarily Oxisols and Ultisols in South America and Oxisols in Africa (Narro et al., 2001). In Brazil, the “cerrado” (tropical savannah) is the main grain-producing region and occupies approximately 205 million ha, which is approximately 23% of the country. The soils in this region are predominantly Oxisols (46%), Ultisols (15%) and Entisols (15%), with low fertility, high aluminium saturation, and high P-fixation capacity (Fageria and Stone, 1999). Low fertility is characteristic of acidic soil; therefore, the correction of these soils is

very important for the proper growth of crops (Soratto and Cruciol, 2008a,b; Soratto et al., 2010).

Liming is the most commonly used practice to neutralize soil acidity and restore production capacity, increase nutrient availability, and reduce levels of toxic elements (Caires et al., 2001; Pavan and Oliveira, 2000). In conventional tillage systems, lime is incorporated into the soil by ploughing and disking. This practice breaks up the soil aggregates, exposes the soil and increases susceptibility to erosion (Bronick and Lal, 2005). Additionally, aggregate disruption promotes the mineralization of previously protected organic matter (Caires et al., 2006a; Westerhof et al., 1999). Therefore, the interest in surface liming to control soil acidity since the implementation of no-till systems is mainly based on preserving the physical properties of the soil (Caires et al., 2011) and maintaining agricultural sustainability in tropical and subtropical regions (Caires et al., 2005).

However, in the short term, the effects of superficial liming are restricted to the soil surface (Pavan and Oliveira, 2000) because without incorporation, there is less contact between the particles

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of lime and soil colloids (Caires et al., 2005). The neutralization of the soil acidity in the subsurface with limestone is generally slow, particularly in variably charged soils (Ernani et al., 2004), because the movement of limestone to greater depths varies according to the time, dose, form, and frequency of application; soil type; climate; acid fertilizer management and crop system (Blevins et al., 1978; Caires et al., 2008b,c, 2005; de Oliveira and Pavan, 1996).

Phosphogypsum is more soluble than limestone and is composed mostly of calcium sulphate ( $\text{CaSO}_4 \cdot \text{H}_2\text{O}$ ); however, its addition does not change the soil pH. The phosphogypsum that is applied to the soil surface moves along the profile under the influence of percolating water (Caires et al., 1999). As a result, there is an increase in the supply of calcium and a reduction in the aluminium toxicity in the subsoil (Caires et al., 2003, 1999). Phosphogypsum is therefore an alternative for improving the root environment in the subsoil and can be used in acidic soils as a supplement for liming (Caires et al., 2003).

In the subtropical region of Brazil, where the rainfall is well distributed throughout the year, there are several reports of a lack of response of crops to the application of soil acidity correctives in no-till systems (Caires et al., 2011, 2008b,c, 2006b; Moreira et al., 2001). These findings have been attributed to the greater accumulation of organic matter and nutrients in the surface that reduce the activity of Al and consequently its toxicity by forming Al–organic complexes and to the increased ionic strength of the soil solution (Nolla and Anghinoni, 2006; Vieira et al., 2009).

It is possible the vast majority of the cerrado biome can experience increases in grain yield and fibre production with lime and phosphogypsum application in no-till systems, even on the surface, unlike what has been observed elsewhere in the subtropical region. This possibility exists because in these regions, the vast majority of cultivated areas have a low organic matter content and a low amount of straw on the surface, leading to low water storage and high evaporation of the soil water, respectively.

Thus, the probable soil acidity correction, aluminium reduction levels and base saturation elevation, especially that of calcium, in the soil profile in a short time will enable greater root depth due to the mechanisms that promote the movement of compounds produced by the dissociation of limestone and phosphogypsum. This condition will increase the water stress tolerance of plants when dry spells induce water stress, especially during off-season cultivation.

The soil acidity correction dynamics in the soil surface in no-till systems and the benefits of the joint application of limestone and phosphogypsum in long-term experiments remain poorly investigated, especially in cerrado conditions in tropical regions. However, this knowledge is important for adjusting the limestone and phosphogypsum recommendations for annual crops in no-till systems.

This study evaluated the changes in soil chemical attributes and the root growth, nutrition and productivity of soybean, black oat and sorghum resulting from limestone and phosphogypsum surface application in an established no-till system in a tropical region.

## 2. Materials and methods

### 2.1. Site description

This experiment was conducted in Botucatu, State of São Paulo, southeastern Brazil, ( $48^{\circ}23'W$ ,  $22^{\circ}51'S$  and 765 masl) during the 2008/2009 and 2009/2010 growing seasons. Soil at this location is classified as Rhodic Ferralsol (FAO, 2006) [kaolinitic, thermic Typic Haplorthox (Soil Survey Staff, 2014), with sandy loam texture] and had been managed since 2002 in a no-till system as follows: growing season 2002/2003, rice in the summer and black oat in the fall; growing season 2003/2004, bean in the summer and black oat in the fall; growing seasons 2004/2005 and 2005/2006, peanut in the

summer and oat in the fall; and growing seasons 2006/2007 and 2007/2008, intercropped corn with *Urochloa ruziziensis*.

The climate according to Köeppen's classification is Cwa (tropical, with a dry winter and hot, rainy summer). Monthly mean values of rainfall and temperature during the experiment are shown in Fig. 1.

At the beginning of the experiment (October 2002) and before lime and phosphogypsum reapplication (August 2004), the chemical attributes of the topsoil (0–0.20 m) were determined (Table 1). In August 2004, soil samples were collected (0–0.20 m) for soil particle size distribution determination with the following results: sand, silt, and clay contents of 540, 110, and  $350 \text{ g kg}^{-1}$ , respectively. In the subsoil (0.20–0.40 m), the clay content was  $360 \text{ g kg}^{-1}$ . The bulk density at depth 0–0.20 m was  $1.128 \text{ Mg m}^{-3}$ .

### 2.2. Experimental design and treatments

The experimental design was a randomized complete block with 4 replications. The experimental units were  $5.4 \times 9 \text{ m}$ . The treatments were as follows: original condition, lime application (Eq. (1), to increase the base saturation to 70%), phosphogypsum application (Eq. (4)), and combined lime (Eq. (1), to increase the base saturation to 70%) plus phosphogypsum (Eq. (4)) application.

At the beginning of the experiment (October 15, 2002), limestone was applied superficially at a rate of  $2700 \text{ kg ha}^{-1}$ . Phosphogypsum was applied one day after liming at a rate of  $2100 \text{ kg ha}^{-1}$ . The reapplication was based on a soil analysis that was carried out in August 2004, where the base saturation in the treatment in which the limestone was applied reached 50%, the pre-established critical level for the reapplication. Thus, on November 19, 2004, the application of limestone and phosphogypsum was performed at rates of  $2000 \text{ kg ha}^{-1}$  and  $2100 \text{ kg ha}^{-1}$ , respectively.

### 2.3. Dolomitic limestone and phosphogypsum characteristics

The dolomitic limestone composition was 23.3% CaO, 17.5% MgO, and 71% effective calcium carbonate equivalence (ECCE). In a physical analysis of the dolomitic limestone, 68.8, 92.4, and 99.7% of the particles passed through 50-, 20-, and 10-mesh sieves, respectively. The composition of phosphogypsum ( $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ ), a by-product obtained from the Brazilian phosphoric acid industry, was 20% Ca, 16% S, and a small residue of 0.1% P and F. In a physical analysis of the phosphogypsum, 60 and 90% of the particles were retained in 50- and 20-mesh sieves, respectively.

The dolomitic limestone rate (LR) was calculated to increase the base saturation in topsoil (0–20 cm) to 70% according to Eq. (1):

$$\text{LR}(\text{kg ha}^{-1}) = \frac{(\text{BS}_2 - \text{BS}_1)(\text{CEC}/\text{BD})}{\text{ECCE}/100} \quad (1)$$

where  $\text{BS}_2$  is the estimated base saturation (70%) and  $\text{BS}_1$  is the base saturation as measured in the soil analysis as in Eq. (2):

$$\text{BS}_1(\%) = \frac{(\text{Ca}_{\text{ex}} + \text{Mg}_{\text{ex}} + \text{K}_{\text{ex}})100}{\text{CEC}} \quad (2)$$

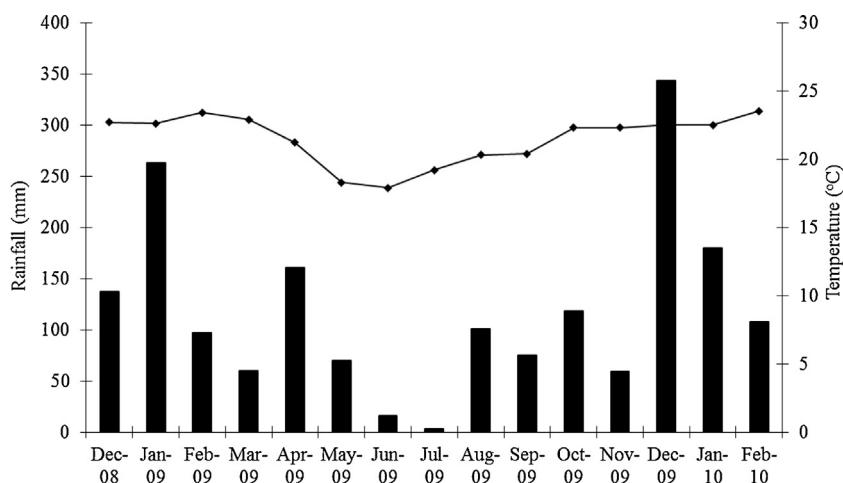
where  $\text{Ca}_{\text{ex}}$ ,  $\text{Mg}_{\text{ex}}$ , and  $\text{K}_{\text{ex}}$  are basic exchangeable cations, BD is bulk density at 0–0.20 m depth, and CEC is the total cation exchange capacity, as calculated by Eq. (3):

$$\begin{aligned} \text{CEC}(\text{cmol}_c \text{ kg}^{-1}) &= \text{Ca}_{\text{ex}} \\ &+ \text{Mg}_{\text{ex}} + \text{K}_{\text{ex}} + \text{total acidity at pH } 7.0 (\text{H} + \text{Al}) \end{aligned} \quad (3)$$

The phosphogypsum rate (GR) was calculated using Eq. (4):

$$\text{GR}(\text{kg ha}^{-1}) = 6\text{CL}(\text{van Raij et al., 1997}) \quad (4)$$

where CL is the clay content ( $\text{g kg}^{-1}$ ) in the soil layer from 20 to 40 cm.



**Fig. 1.** Monthly rainfall (mm) and average temperature (°C) in the experimental area at Botucatu, São Paulo State, Brazil, during the period from December 2008 to February 2010.

**Table 1**

Chemical characteristics of the soil before the experiment (October 2002) and before the last application (August 2004).

Depth (m)	pH (CaCl <sub>2</sub> )	SOM (g kg <sup>-1</sup> )	P (resin) (mg kg <sup>-1</sup> )	H + Al (mmol <sub>c</sub> kg <sup>-1</sup> )	Al (mmol <sub>c</sub> kg <sup>-1</sup> )	K (mmol <sub>c</sub> kg <sup>-1</sup> )	Ca (mmol <sub>c</sub> kg <sup>-1</sup> )	Mg (mmol <sub>c</sub> kg <sup>-1</sup> )	CEC (mmol <sub>c</sub> kg <sup>-1</sup> )	BS (%)
October 2002										
0–0.20	4.2	24	10	42	7.3	1.4	16	6	65	37
August 2004										
0–0.20	4.9	30	39	39	2.6	1.2	27	11.0	79	50

#### 2.4. Crop characteristics and management

Soybean (*Glycine max*) was sown on December 11, 2008, and October 29, 2009, using the cultivars MGBR-46 and CD 216, respectively. The row spacing was 0.45 m with approximately 22 and 20 seeds m<sup>-1</sup>, respectively. The seeds were treated with a fungicide (Vitavax + thiram – 50 + 50 g of the active ingredient for every 100 kg of seed) and an inoculant (*Bradyrhizobium japonicum*). The base fertilization consisted of 250 kg ha<sup>-1</sup> of 04–20–20 NPK formula according to the results of a soil chemical analysis and soybean crop recommendations (van Raij et al., 1997). In both years, phytosanitary measures were taken based on soybean recommendations and requirements.

Black oat (*Avena strigosa*) and sorghum (*Sorghum bicolor* (L.) Moench) were sown on June 22, 2009, and March 08, 2010, using the cultivar “Comum” and the hybrid AG 1020, respectively. The row spacing was 0.17 and 0.60 m, respectively, with seed rates of 60 kg ha<sup>-1</sup> and 18 seeds m<sup>-1</sup>, respectively. The seeds were treated with a fungicide (Vitavax + thiram – 50 + 50 g of the active ingredient for every 100 kg of seed). In both crops, no fertilizer was applied, and phytosanitary measures were taken based on recommendations and requirements.

Soybeans were harvested on April 24, 2009, and February 24, 2010; black oat and sorghum were harvested on October 19, 2009, and July 21, 2010, respectively. Samples were taken for the evaluation of the yield components and final yield (130 g kg<sup>-1</sup> moisture content). The yield components were plant population, number of pods per plant, number of grains per pod and 100-grain weight for soybean, and panicles per m<sup>2</sup>, grains per panicle, and 1000-grain weight for black oat and sorghum.

#### 2.5. Soil sampling and chemical analysis

Soil samples were taken 48 and 60 months after lime and phosphogypsum reapplication at depths of 0–0.05, 0.05–0.10,

0.10–0.20, 0.20–0.40, and 0.40–0.60 m. Seven subsamples were taken randomly from the plot area of each of the plots between rows and combined into one composite sample. Sample analyses included the following. The soil pH was determined in a 0.01 mol L<sup>-1</sup> CaCl<sub>2</sub> suspension (1:2.5 soil:solution). The soil organic matter was determined by the Walkley–Black method (Walkley and Black, 1934). The total acidity at pH 7.0 (H + Al) was estimated by the SMP-buffer solution method (van Raij et al., 2001). The exchangeable Al was extracted with neutral 1 mol L<sup>-1</sup> KCl at a 1:10 soil:solution ratio and determined by titration with a 0.025 mol L<sup>-1</sup> NaOH solution. Phosphorus and exchangeable Ca, Mg and K were extracted with ion-exchange resins; in the extract, P was determined colorimetrically, and cations by atomic absorption spectrometry (van Raij et al., 2001). Using the exchangeable bases and total acidity at pH 7.0 (H + Al) results, the base saturation values were calculated (van Raij et al., 2001). The soil SO<sub>4</sub>-S analyses were performed through extraction by calcium phosphate 0.01 mol L<sup>-1</sup> at a 1:2.5 soil:solution ratio and later determined by the turbidimetric method by precipitating sulphate as BaSO<sub>4</sub> (Vitti, 1988).

#### 2.6. Soybean root sampling and analyses

The soybean root system was sampled when 50% of the plants of each subplot were flowering. A 4.5-cm diameter galvanized-steel auger was used for sampling at depths of 0–0.05, 0.05–0.10, 0.10–0.20, 0.20–0.40, and 0.40–0.60 m. For each layer, five subsamples were collected from under the soybean plant rows and five in the middle of the inter-rows (0.20 m from the plant row) of each plot. The subsamples (within rows and between rows) were combined into a composite sample. The root samples were washed with tap water through a 0.5-mm mesh sieve to avoid the excessive loss of material; the debris was removed by hand. Then, the roots were immersed in a 30% ethyl alcohol solution in plastic pots with lids and stored under refrigeration at 2 °C until evaluation. To evaluate

**Table 2**

Probabilities of the calculated *F*-value for soil chemical properties at the depths of 0–0.05, 0.05–0.10, 0.10–0.20, 0.20–0.40 and 0.40–0.60 m at 48 and 60 months after reapplication of lime and phosphogypsum.

ate the root length density ( $\text{km m}^{-3}$ ) and diameter (cm), the roots were digitalized using an optical scanner (Scanjet 4C/T, HP-USA, Palo Alto, CA) with a resolution of 250 dpi and analysed using "Win Rhizo" version 3.8-b (Regent Instrument Inc., Quebec, Canada).

## 2.7. Sampling and tissues analyses

Soybean, black oat and sorghum leaf sampling for nutrient concentration analysis was performed when 50% of the plants were in the full flowering stage. The sampling was randomized by choosing 30 plants per plot and sampling the 3rd leaf with the petiole for soybean (Ambrosano et al., 1997), the fourth leaf sheath visible counted from the apex for sorghum, and the flag leaf for black oat (Cantarella et al., 1997). At this stage, ten plants were sampled to determine the dry matter, which was expressed in  $\text{kg ha}^{-1}$ . The plant and leaf samples were dried by forced-air circulation at 65 °C for 72 h, and then the leaves were ground. The concentrations of N, P, K, Ca, Mg, S, Cu, Zn, Fe and Mn in the leaves were determined according to the methods described by Malavolta et al. (1997). Nitrogen was digested with  $\text{H}_2\text{SO}_4$ , and the other nutrients were digested with a nitro-perchloric solution. From the digested solution, N, P and S were determined by the Kjeldahl distillation, colorimetry and turbidimetry methods, respectively, and, K, Ca, Mg, Cu, Zn, Fe and Mn were determined by atomic absorption spectrophotometry.

## 2.8. Statistical analysis

For all of the data, a two-way ANOVA was run using the statistical software package SISVAR (UFLA—Lavras, Minas Gerais, Brazil) (Ferreira, 2008). The blocks and all of the block interactions were considered random effects. The treatments and growing seasons were considered fixed effects. For the significant effects, the means were separated using Fisher's protected LSD test. For the soil chemical characteristics, nutrient, aboveground biomass, yield components and grain yield data, the effects were considered significant at  $P \leq 0.05$ , whereas the effects were considered significant at  $P \leq 0.10$  for the root characteristic data. According to Gregory (2006), 15–20 root samples are necessary to detect significant differences at the 10% level of probability, whereas significance at the 5% level often requires approximately 60–90 samples in each experimental unit. Therefore, a higher probability for the root system evaluation was used because these characteristics are more variable (Fageria and Moreira, 2011).

## 3. Results

### 3.1. Soil chemical properties

Surface liming increased the soil pH in the surface layers after 48 months to a depth of 0.20 m, extending throughout the soil profile after 60 months of reapplication (Table 2 and Fig. 2). Phosphogypsum also influenced the soil pH, with higher values than the control, reaching the same depths as observed with liming.

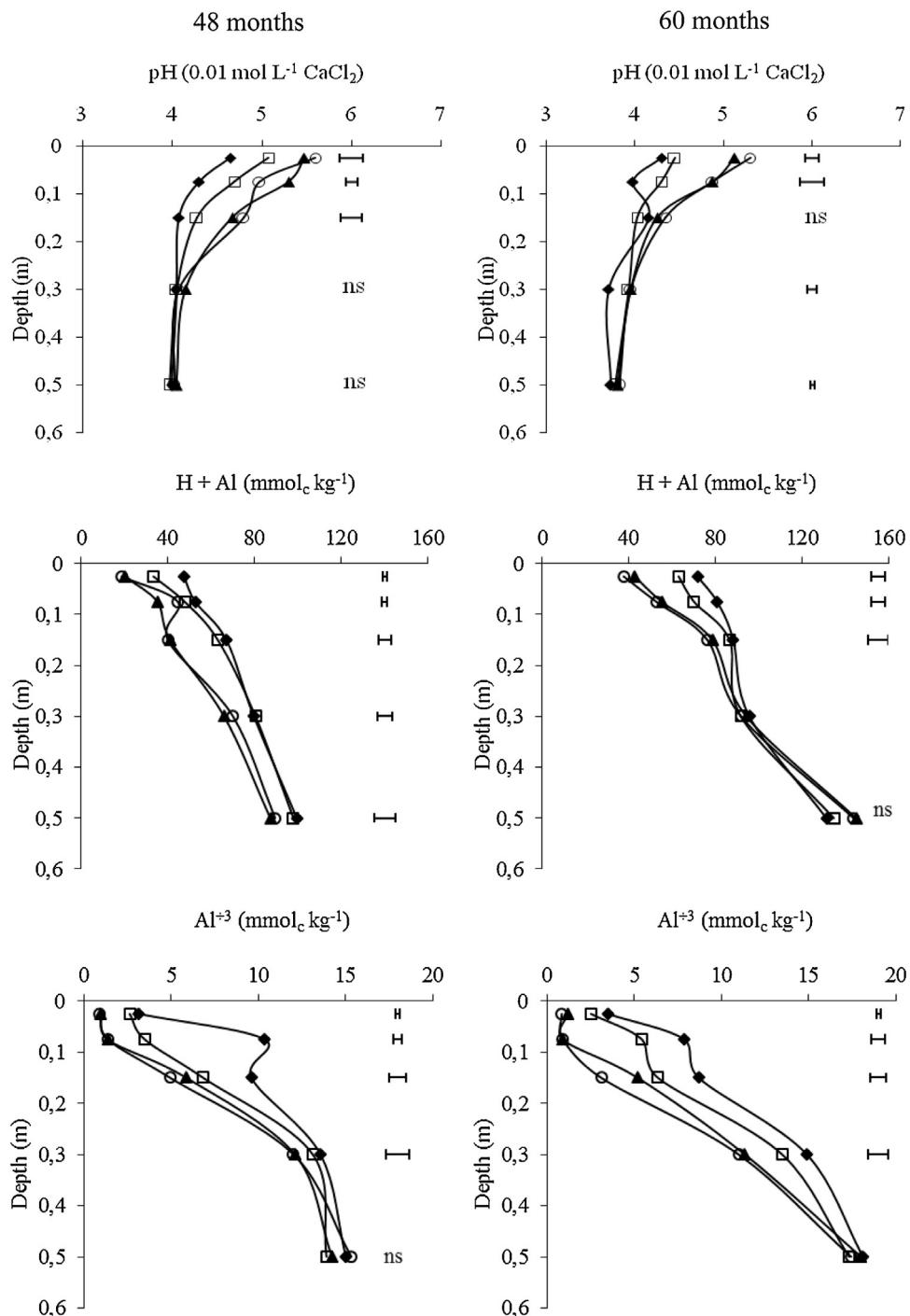
After 48 months, liming reduced the H + Al contents to a depth of 0.60 m, and after 60 months, H + Al was restricted to a depth of 0–0.20 m (Table 2 and Fig. 2). The phosphogypsum also reduced the H + Al contents compared to the control to a depth of 0.10 m at 48 months and persisted until 60 months after reapplication. Al<sup>3+</sup> levels were reduced to a depth of 0.40 m with liming and phosphogypsum after 48 months and persisted until 60 months after reapplication (Table 2 and Fig. 2).

Phosphogypsum when applied with liming contributed little to improving the soil acidity components in the long term. With respect to soil acidity, 60 months after reapplication, there was

Table 3

Factor	Root length density ( $\text{km m}^{-3}$ )					Root diameter (cm)
	0–0.05 m	0.05–0.10 m	0.10–0.20 m	0.20–0.40 m	0.40–0.60 m	
Control	122	75	35	9	4	0.050
Gypsum (G)	120	78	38	8	6	0.047
Lime (L)	114	71	24	6	5	0.046
L + G	112	78	23	9	8	0.047
Growing season						
2008/2009	103 b <sup>a</sup>	70 b	34 a	10 a	7 b	0.047
2009/2010	130 a	82 a	26 b	6 b	5 b	0.048
Block	0.950	0.221	0.919	0.349	0.357	0.689
Treatments (T)	0.113	0.356	0.217	0.240	0.334	0.119
Growing season (Y)	0.002	0.001	0.047	<0.001	0.275	0.155
T × Y	0.157	0.202	0.749	0.128	0.477	0.572
						0.172
						0.315
						0.201
						0.1172

<sup>a</sup> Means followed by different letters in the column differ statistically by the LSD test ( $P < 0.05$ ).



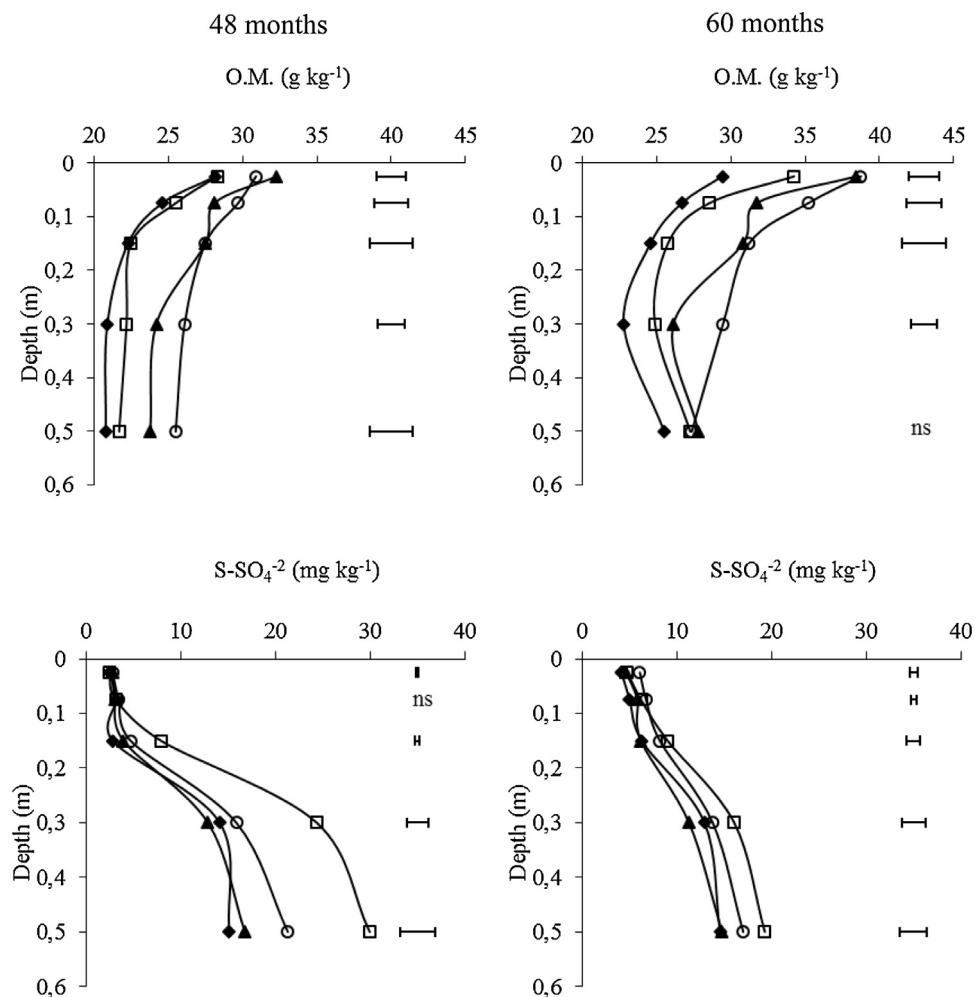
**Fig. 2.** pH, H + Al and Al<sup>3+</sup> levels 48 and 60 months after the reapplication of lime (▲), phosphogypsum (□), and limestone + phosphogypsum (○) and the corresponding levels of the control (◆). Horizontal bars indicate the least significant difference (LSD).

amelioration in the components of acidity, indicating that most of the applied products had already reacted.

The soil organic matter increased with liming (Table 2 and Fig. 3) throughout the soil profile 48 months after reapplication, and this effect was observed up to 0.40 m 60 months after reapplication. Liming, when associated with phosphogypsum, provided the highest levels in the deeper layers (0.20–0.60 m) after 48 months; after 60 months, it provided the highest levels only in the 0.05–0.10-m and 0.20–0.40-m layers. The effects of phosphogypsum were only observed after 60 months at depths of 0–0.05 and 0.20–0.40 m compared to the control.

The S-SO<sub>4</sub><sup>2-</sup> levels increased in the soil layers of 0.10–0.60 m with phosphogypsum application, with or without liming, after 48 months (Table 2 and Fig. 3). After 60 months, this effect was extended to the layer of 0.05–0.60 m. In the most superficial layer of the soil (0–0.05 m), only an effect of phosphogypsum was observed.

Liming after 48 months increased K levels at a depth of 0.05–0.20 m and P at a depth of 0–0.20 m, and when associated with phosphogypsum, the P levels were even higher in the layer from 0 to 0.10 m (Table 2 and Fig. 4). After 60 months, the effects of lime, with or without phosphogypsum, on K levels were observed at a depth



**Fig. 3.** Organic matter (O.M.) and S-SO<sub>4</sub><sup>2-</sup> levels 48 and 60 months after the reapplication of lime (▲), phosphogypsum (□), and limestone + phosphogypsum (○) and the corresponding levels of the control (◆). Horizontal bars indicate the least significant difference (LSD).

of 0–0.05 m. However, phosphogypsum application associated with lime no longer changed the P levels after 60 months.

After 48 to 60 months, the Ca<sup>2+</sup> and Mg<sup>2+</sup> levels increased at all depths with liming (Table 2 and Fig. 5). However, when combined with phosphogypsum, the concentrations of Ca<sup>2+</sup> increased throughout the soil profile except in the layer from 0.05 to 0.10 m after 60 months, whereas the Mg<sup>2+</sup> levels decreased at depths of 0.10–0.20 and 0.40–0.60 m after 48 months and in the layer from 0.05 to 0.60 m after 60 months. Phosphogypsum also positively influenced the Ca<sup>2+</sup> levels at all of the depths compared with the control but reduced the Mg<sup>2+</sup> levels at depths of 0.10–0.20 and 0.20–0.40 m after 48 and 60 months, respectively.

The soil amendment effects on the H+Al contents, Ca and Mg were reflected in the base saturation (Table 2 and Fig. 5). Liming and phosphogypsum increased the base saturation throughout the soil profile after 48 and 60 months.

### 3.2. Soybean root growth

The soybean root characteristics were influenced only by the year effect, with no significant interaction (Table 3). Thus, phosphogypsum and lime application did not alter the length or average diameter of the soybean roots. Regarding the year effect, the root growth at the soil surface (0–0.10 m) was higher in 2009/2010, with approximately 85% of the roots found in this layer at this time, whereas in 2008/2009, this proportion was 77%. However,

below a depth of 0.10 m, the highest growth and root diameter were observed in 2008/09.

### 3.3. Nutrition and grain yield

There was no significant interaction between the treatments and year for any evaluated nutrient (Table 4). The P and Fe concentrations in the soybean leaves were not influenced by the treatments, nor were the P, Ca, Cu and Mn concentrations influenced by the year. The N, Ca, Mg and S concentrations increased with liming and were even higher when associated with phosphogypsum with the exception of Mg. However, the micronutrient concentrations (Cu, Zn and Mn) were inversely proportional to the macronutrient concentrations. Regarding the effect of year, the highest concentrations of N, K, Mg, S, Cu, Zn and Fe were found in 2008/2009.

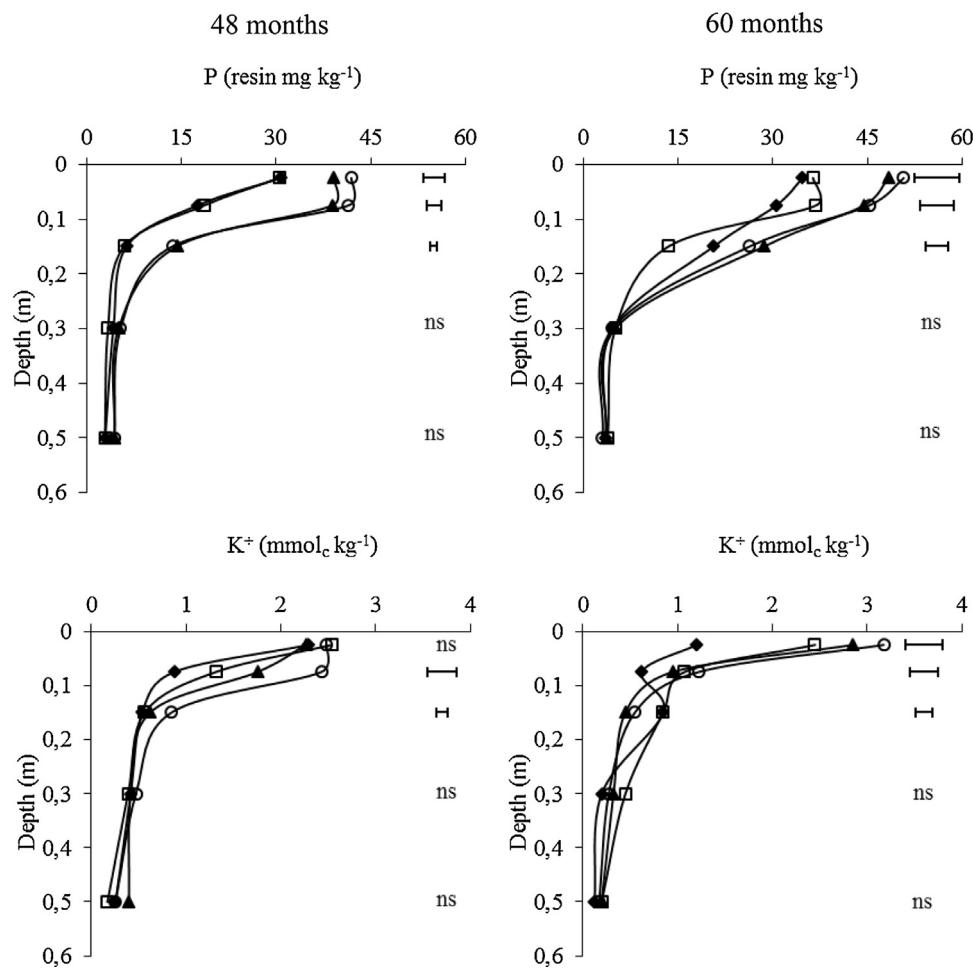
The yield components and shoot dry matter of soybean were influenced by the effects of the treatments and year; however, there were no significant interactions (Table 4). The surface application of lime increased the shoot dry matter, plant population and the 100-grain weight, and when associated with phosphogypsum, the number of pods per plant was higher, directly reflecting the grain yield. The application of limestone alone and in association with phosphogypsum increased the grain yield by 17 and 29%, respectively. In 2008/2009, the production of shoot dry matter, plant population and 100-grain weight were higher; however, the aver-

**Table 4**

Nutrient concentrations (N, P, K, Ca, Mg, S, Cu, Zn, Mn and Fe), shoot dry matter production, yield components (plant population, pods per plant, grains per pod, and 100-grain weight) and soybean yield affected by surface liming, phosphogypsum and lime + phosphogypsum application under a no-tillage system. Botucatu, State of São Paulo, Brazil, 2008–2010.

Factor	N (g kg <sup>-1</sup> )	P (g kg <sup>-1</sup> )	K (g kg <sup>-1</sup> )	Ca (g kg <sup>-1</sup> )	Mg (g kg <sup>-1</sup> )	S (g kg <sup>-1</sup> )	Cu (mg kg <sup>-1</sup> )	Zn (mg kg <sup>-1</sup> )	Mn (mg kg <sup>-1</sup> )	Fe (mg kg <sup>-1</sup> )
<b>Treatments</b>										
Control	40.6 c <sup>a</sup>	4.7 a	28.1 b	17.8 d	6.0 c	2.5 c	18 a	94 a	194 a	200 a
Gypsum	42.9 b	4.8 a	30.4 a	20.5 c	6.0 c	2.7 b	21 a	85 b	159 b	213 a
Lime	43.5 b	4.6 a	27.6 b	21.9 b	6.7 a	2.8 ab	12 b	67 c	94 c	199 a
Lime + gypsum	45.6 a	4.7 a	24.4 c	23.5 a	6.4 b	3.0 a	8 c	61 d	88 c	200 a
<b>Growing season</b>										
2008/2009	44.2 a	4.8 a	30.0 a	20.6 a	6.5 a	3.1 a	16 a	79 a	138 a	226 a
2009/2010	42.1 b	4.6 a	25.3 b	21.2 a	6.1 b	2.4 b	14 a	74 b	129 a	180 b
<i>F</i> probability										
Block	0.125	0.992	0.999	0.080	0.281	0.912	0.846	0.256	0.300	0.900
Treatments (T)	<0.001	0.772	<0.001	<0.001	<0.001	<0.001	<0.001	<0.001	<0.001	0.108
Growing season (Y)	<0.001	0.212	<0.001	0.171	<0.001	<0.001	0.304	0.015	0.460	<0.001
T × Y	0.135	0.470	0.285	0.257	0.173	<0.001	0.203	0.300	0.979	0.099
Dry matter (kg ha <sup>-1</sup> )		Population (pl ha <sup>-1</sup> )		Pods per plant (number)		Grains per pod (number)		100-grain weight (g)		Grain yield (kg ha <sup>-1</sup> )
<b>Treatments</b>										
Control	3.144 b <sup>a</sup>		222 b		35.1 b		2.25 a		15.90 b	
Gypsum	3.205 b		233 ab		36.8 b		2.30 a		15.95 b	
Lime	4.173 a		251 a		37.3 ab		2.30 a		17.43 a	
Lime + gypsum	4.359 a		250 a		39.6 a		2.35 a		17.25 a	
<b>Growing season</b>										
2008/2009	4.037 a		262 a		36.3 a		1.85 b		17.92 a	
2009/2010	3.404 b		216 b		38.1 a		2.75 a		15.34 b	
<i>F</i> probability										
Block	0.995		0.949		0.917		0.063		0.111	
Treatments (T)	<0.001		<0.001		0.007		0.173		<0.001	
Growing season (Y)	<0.001		<0.001		0.050		<0.001		<0.001	
T × Y	0.112		0.139		0.525		0.731		0.129	

<sup>a</sup> Means followed by different letters in the column differ statistically by the LSD test (*P*<0.05).



**Fig. 4.** P and K levels 48 and 60 months after the reapplication of lime (▲), phosphogypsum (□), and limestone + phosphogypsum (○) and the corresponding levels of the control (◆). Horizontal bars indicate the least significant difference (LSD).

age number of grains per pod was higher in 2009/2010 and reflected the higher grain yield this season.

In the black oat crop, the N, P, K, Cu and Zn concentrations were not affected by product application (Table 5). However, the Ca concentration increased with product application, with the highest concentrations observed with liming. The Mg concentration increased only with liming, whereas that of S increased with phosphogypsum application, with or without liming. The Mn and Fe concentrations decreased with liming, and the decrease in Mn was greater when liming was associated with phosphogypsum.

Regarding the shoot dry matter and yield components of black oat, liming increased the shoot dry matter and number of panicles per m<sup>2</sup> (Table 5). However, the number of panicles per m<sup>2</sup> was lower when liming was combined with phosphogypsum, reducing the yield by 6%; however, compared with the control, the yields with lime alone and lime associated with phosphogypsum were higher, at 54 and 44%, respectively.

Liming increased the uptake of N, Ca, Mg and S by sorghum and that of Ca was greater when liming was combined with phosphogypsum (Table 6). Phosphogypsum application only influenced the absorption of S compared to that of the control. Regarding micronutrients, liming, with or without phosphogypsum only reduced the uptake of Mn.

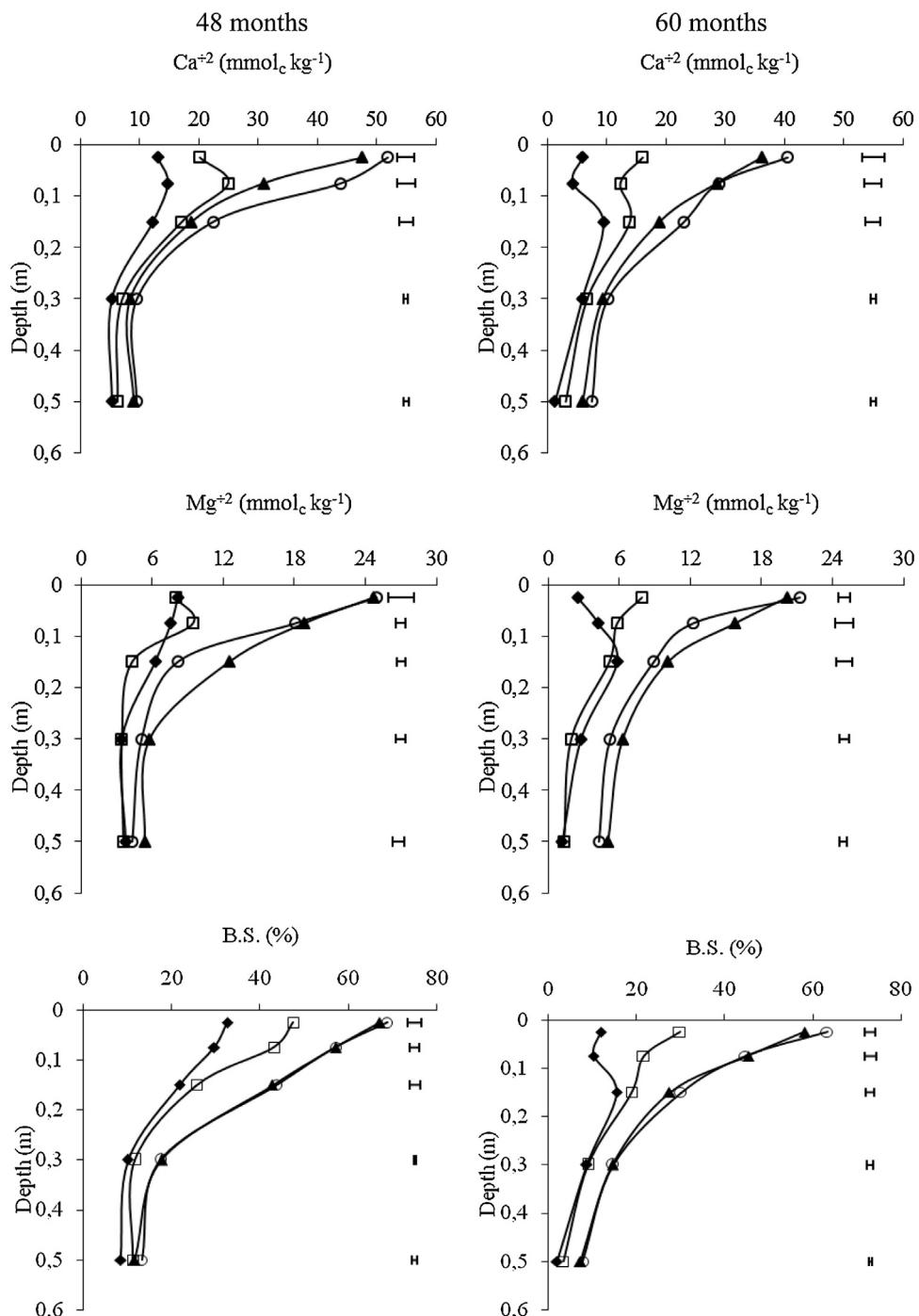
The shoot dry matter, number of panicles per m<sup>2</sup>, and number of grains per panicle of sorghum were influenced by lime and phosphogypsum application, which was reflected directly in the grain yield (Table 6). Thus, the grain yield with the application of phosphogypsum, limestone and limestone associated with phosphogypsum

was 191, 1406 and 1589 kg ha<sup>-1</sup> higher, respectively, than that of the control. We note that each crop has a different response to soil amendments, which may be related to intrinsic characteristics, such as the weather conditions during development.

## 4. Discussion

### 4.1. Soil chemical properties

Surface lime application corrected the soil acidity in the soil profile (Fig. 2). Caires et al. (2005) and Caires et al. (2011) found that surface lime application under no-till in southern Brazil, a subtropical region, reduced the soil acidity in the long term throughout the soil profile after 120 months in a loamy Typic Hapludox (0–0.20 m) and after 84 months in a clayey Rhodic Hapludox (0–0.60 m), respectively. Acidity correction occurs with an increase in the concentration of hydroxyl, a limestone dissociation product in the soil, as a fraction of this hydroxyl group reacts with excess H<sup>+</sup> in the solution, reducing the concentration of H<sup>+</sup>/Al and increasing the soil pH, while the remainder promotes aluminium precipitation in the form of Al(OH)<sub>3</sub>, an element which is not toxic to plants (de Oliveira and Pavan, 1996). Thus, subsoil correction can occur due to several factors, such as the physical displacement of the fine particles of limestone through the water movement via channels that are formed by dead roots and organisms (de Oliveira and Pavan, 1996; Rheinheimer et al., 2000) and through weakness planes (Amaral et al., 2004), which are kept intact due the minimum soil disturbance in a no-till system. Another factor that can be attributed to



**Fig. 5.** Ca and Mg levels and Base saturation (B.S.) 48 and 60 months after the reapplication of lime (▲), phosphogypsum (□), and lime + phosphogypsum (○) and the corresponding levels of the control (◆). Horizontal bars indicate the least significant difference (LSD).

the effect of surface liming in the subsoil is the movement of ions. The increased soil pH on the soil surface due to superficial application can accelerate the rate at which  $\text{HCO}_3^-$  ions, followed by Ca and Mg, move into the subsoil to react to acidity (Caires et al., 2003). According to Rheinheimer et al. (2000), the effects of liming at depth only occurred when the pH in the water in the limestone dissolution zone was between 5.2 and 5.6. In this situation, the formation and migration of  $\text{Ca}(\text{HCO}_3)_2$  and  $\text{Mg}(\text{HCO}_3)_2$  occurred in the subsurface layers. This effect with depth is known as the “alkalizing front”, as noted by several authors (Caires et al., 2008c; Soratto and Cruciol, 2008b).

Surface phosphogypsum application, compared to the control, influenced the soil acidity components (Fig. 2). Caires et al. (2003) observed an increase in the soil pH in the layers of 0.20–0.40 (8 months) and 0.40–0.60 m (8, 20 and 32 months), and Soratto and Cruciol (2008b) reported a decrease in the exchangeable Al at all depths 12 months after application. These results have been attributed to an exchange reaction of ligands on the surface of soil particles involving hydrated oxides of iron and aluminium, with  $\text{SO}_4^{2-}$  displacing  $\text{OH}^-$  and thus promoting the partial neutralization of acidity (Reeve and Sumner, 1972). In addition, aluminium precipitation could also occur with the formation of minerals (Adams and Rawajfih, 1977). Another possibility is the leaching of

**Table 5**

Nutrient concentrations (N, P, K, Ca, Mg, S, Cu, Zn, Mn and Fe), shoot dry matter production, yield components (panicles per square metre, grains per panicle, 1000-grain weight and spikelet fertility) and black oat yield affected by surface liming, phosphogypsum and lime + phosphogypsum application under a no-tillage system. Botucatu, State of São Paulo, Brazil, 2008–2010.

	N (g kg <sup>-1</sup> )	P (g kg <sup>-1</sup> )	K (g kg <sup>-1</sup> )	Ca (g kg <sup>-1</sup> )	Mg (g kg <sup>-1</sup> )	S (g kg <sup>-1</sup> )	Cu (mg kg <sup>-1</sup> )	Zn (mg kg <sup>-1</sup> )	Mn (mg kg <sup>-1</sup> )	Fe (mg kg <sup>-1</sup> )								
<b>Treatments</b>																		
Control	19.7 a <sup>a</sup>	3.9 a	30.9 a	10.6 c	4.6 b	2.5 bc	10.4 a	55 a	259 a	173 a								
Gypsum (G)	20.5 a	3.9 a	28.3 a	11.6 b	4.7 b	2.9 a	10.9 a	59 a	268 a	177 a								
Lime (L)	19.4 a	3.9 a	29.5 a	13.0 a	5.4 a	2.4 c	10.4 a	53 a	195 b	115 b								
L + G	19.1 a	3.7 a	31.7 a	12.6 a	4.9 b	2.7 ab	10.9 a	50 a	147 c	124 b								
<i>F</i> probability																		
Block	0.999	0.895	0.555	0.250	0.431	0.362	0.146	0.136	0.354	0.265								
Treatments	0.726	0.423	0.248	<0.001	<0.001	0.007	0.889	0.159	<0.001	0.001								
Dry matter (kg ha <sup>-1</sup> )		Panicles m <sup>-2</sup> (number)		Grains per panicle (number)		1000-grain weight (g)		Spikelet fertility (%)		Grain yield (kg ha <sup>-1</sup> )								
<b>Treatments</b>																		
Control	4632 c		138 c		46 a		18.6 a		84 a	960 c								
Gypsum (G)	4868 c		147 c		45 a		18.9 a		83 a	1009 c								
Lime (L)	5856 a		216 a		46 a		19.1 a		80 a	1476 a								
L + G	5484 a		205 b		45 a		19.4 a		79 a	1382 b								
<i>F</i> probability																		
Block	0.215		0.252		0.101		0.835		0.118	0.224								
Treatments	<0.001		<0.001		0.943		0.761		0.367	<0.001								

<sup>a</sup> Means followed by different letters in the column differ statistically by the LSD test (*P*<0.05).

**Table 6**

Nutrient concentrations (N, P, K, Ca, Mg, S, Cu, Zn, Mn and Fe), shoot dry matter production, yield components (panicles per square metre, grains per panicle and 1000-grain weight) and sorghum yield affected by surface liming, phosphogypsum and lime + phosphogypsum application under a no-tillage system. Botucatu, State of São Paulo, Brazil, 2008–2010.

	N (g kg <sup>-1</sup> )	P (g kg <sup>-1</sup> )	K (g kg <sup>-1</sup> )	Ca (g kg <sup>-1</sup> )	Mg (g kg <sup>-1</sup> )	S (g kg <sup>-1</sup> )	Cu (mg kg <sup>-1</sup> )	Zn (mg kg <sup>-1</sup> )	Mn (mg kg <sup>-1</sup> )	Fe (mg kg <sup>-1</sup> )									
<b>Treatments</b>																			
Control	27.2 b <sup>a</sup>	3.9 a	18.0 a	9.4 c	3.8 b	1.8 c	9.3 a	47 a	42 a	344 a									
Gypsum (G)	27.8 b	4.0 a	17.6 a	10.1 c	3.9 b	2.1 b	8.8 a	45 a	39 a	323 a									
Lime (L)	28.8 ab	4.0 a	17.9 a	11.2 b	4.1 ab	2.7 a	6.8 a	43 a	35 b	314 a									
L + G	30.1 a	3.9 a	18.7 a	12.7 a	4.7 a	2.7 a	8.0 a	39 a	32 b	297 a									
<i>F</i> probability																			
Block	0.301	0.113	0.703	0.985	0.912	0.201	0.062	0.827	0.771	0.091									
Treatments	0.012	0.080	0.171	<0.001	0.039	<0.001	0.118	0.369	<0.001	0.616									
Dry matter (kg ha <sup>-1</sup> )		Panicles m <sup>-2</sup> (number)		Grains per panicle (number)		1000-grain weight (g)		Grain yield (kg ha <sup>-1</sup> )											
<b>Treatments</b>																			
Control	1513 d		11.0 d		11 c		36.6 a		45 d										
Gypsum (G)	2307 c		11.9 c		57 b		34.1 a		236 c										
Lime (L)	4084 b		17.2 b		245 a		34.7 a		1451 b										
L + G	6750 a		18.5 a		237 a		35.6 a		1634 a										
<i>F</i> probability																			
Block	0.156		0.899		0.391		0.923		0.185										
Treatments	<0.001		<0.001		<0.001		0.257		<0.001										

<sup>a</sup> Means followed by different letters in the column differ statistically by the LSD test (*P*<0.05).

Al accompanying phosphogypsum, which can be partly favoured by the formation, primarily, of ion pairs or complexes of AlSO<sub>4</sub><sup>+</sup> (Pavan et al., 1984).

In addition, in contrast to the previous soil evaluation, soil acidification occurred (Fig. 2). This effect may suggest that large amounts of lime have already reacted by 60 months because with the total reaction of lime, it is probable that the acidification processes manifest through the pH values, considering the soil buffering capacity, the nutrient extraction processes and the increase in the use of ammonium base nitrogen fertilizer (Caires et al., 2005; Juo et al., 1995). Ciotta et al. (2002) reported soil acidification in a no-till system and explained this effect by the ammonium nitrification process from the acid reaction of fertilizers, which concentrate on the soil surface.

The soil organic matter increased with liming after 48 and 60 months, with or without phosphogypsum (Fig. 3). However, the effects of phosphogypsum were only observed at 60 months. With liming, there is a pH elevation and a concomitant increase in micro-

bial activity (Ekenler and Tabatabai, 2003). This increase accelerates the decomposition of soil organic matter and reduces the stock of soil C (Chan and Heenan, 1999). This effect tends to last only for the short term; however, long-term liming can promote the accumulation of organic matter (Briedis et al., 2012; Hati et al., 2008). This C accumulation in the long-term is related to a greater biomass production per area as a result of soil fertility improvements, such as increased pH, the supply of Ca and Mg, and Al<sup>3+</sup> reduction. It should be taken into account that sixteen crops were grown in sequence in this area in the period between 2002 and 2009 with a significant accumulation of crop residues, including roots, whose development was also improved by soil amendment.

The sulphate levels were influenced only by the application of phosphogypsum, with or without liming, below the surface layer, with reduced levels observed during the study period (Fig. 3). Caires et al. (2006b), in a clayey Rhodic Hapludox, verified an increase in sulphate levels at depths of up to 0.80 m with superficial phosphogypsum application (3, 6 and 9 t ha<sup>-1</sup>) after 53 months, but

only a small fraction was retained in the layer of 0–0.20 m, and the remainder leached into deeper layers. [Caires et al. \(2011\)](#), in a clayey Rhodic Hapludox, also found similar effects with superficial phosphogypsum application ( $9 \text{ Mg ha}^{-1}$ ) after 84 months.

The small retention of sulphate in the upper soil layers is due to the higher pH values and organic matter levels in these layers. The increase in soil pH promotes a predominantly negative net charge, and the organic matter generates a large number of negative charges in the soil, favouring the movement of sulphate ([Camargo and van Raij, 1989](#); [Quaggio et al., 1993](#)). In addition, the sulphate adsorbed on hydroxylated surfaces by specific adsorption decreases with increasing pH ([Rietra et al., 1999](#)).

Surface liming increased the K and P levels in the soil surface, which were intensified by phosphogypsum application after 48 months. After 60 months, the effects of liming persisted for P and exchangeable K when combined with phosphogypsum ([Fig. 4](#)). Several authors have reported that liming can reduce potassium loss by leaching ([Caires et al., 1998](#); [Quaggio et al., 1993, 1982](#)). This effect may be related to increased pH-dependent negative charges caused by liming ([Quaggio et al., 1982](#)) and the charge alteration by divalent cations (Ca and Mg) by forming complexes with water-soluble organic ligands that are present in crop residues ([Miyazawa et al., 1993](#)). In such a scenario, the free charge would be occupied by  $\text{K}^+$ , increasing the exchangeable K levels mainly in the upper soil layers ([Caires et al., 1998](#)).

The increase in the soil pH, increasing the concentration and activity of OH-ions in solution, promotes Fe and Al precipitation. In contrast, the precipitation of low soluble Fe- and Al-phosphates is decreased. In addition, negative charges are generated by hydroxyl deprotonation exposed in clays and organic matter, causing repulsion between the phosphate and the adsorbent surface ([Macbride, 1994](#)). [Fageria and Baligar \(2008\)](#) reported in Brazilian Oxisols a linear increase in available P with increased soil pH in the range of 5.3–6.9, indicating that the increased availability of P in this interval was associated with the release of P ions from Al and Fe oxides, which are responsible for fixing this element in tropical soils. Thus, the annual phosphorus addition by fertilizers provided higher P availability by reducing the force with which P was retained in the soil ([Alvarado and Cajuste, 1993](#)). The increased P availability with phosphogypsum may be related to the competitive adsorption of sulphate and phosphate in acidic conditions ([Geelhoed et al., 1997](#)); furthermore, phosphogypsum composition includes a small content of P.

The exchangeable Ca increased throughout the soil profile and to a greater extent in the upper soil layers with liming and phosphogypsum application ([Fig. 5](#)). Many studies have reported an increase in the soil exchangeable Ca under the influence of liming ([Caires et al., 2011, 2004](#); [Castro and Cruciol, 2013](#); [de Oliveira and Pavan, 1996](#); [Rheinheimer et al., 2000](#); [Soratto and Cruciol, 2008b](#)) or phosphogypsum ([de Oliveira and Pavan, 1996](#); [Rheinheimer et al., 2000,b](#); [Soratto and Cruciol, 2008a,b](#)). Because the two materials are sources of Ca, we observed the effects of lime and phosphogypsum in the present study, with surface application contributing to increased calcium in the soil profile, confirming the results obtained by [Caires et al. \(2011, 2003\)](#) and [Soratto and Cruciol \(2008b\)](#).

Over time, there was a decrease in the exchangeable Ca, similar to the effect that was observed for soil pH values ([Fig. 2](#)), which can be attributed to the extraction of Ca by previous crops. [Caires et al. \(2000\)](#) and [Soratto and Cruciol \(2008b\)](#) also found a reduction in exchangeable calcium levels over time, especially in the surface layers.

Liming increased the exchangeable Mg contents in the soil profile, whereas phosphogypsum application, with or without liming, reduced the exchangeable Mg contents ([Fig. 5](#)). Because dolomitic lime is also an Mg source, the observed increase in the Mg content in the soil profile was expected. Magnesium leaching has frequently

been observed in studies with the surface application of phosphogypsum ([Caires et al., 2006b, 2004, 2003, 1998](#); [Gatiboni et al., 2003](#); [Soratto and Cruciol, 2008b](#)). This movement occurs via the association of cations with the  $\text{SO}_4^{2-}$  anion, forming soluble ions pairs with a lower valence or neutral charge. In this way, these anions have a higher mobility and are more easily leached into the soil profile ([Silva et al., 1998](#)). However, in the present study, the use of dolomitic limestone maintained adequate levels of Mg in the soil for a long period, even when it was associated with phosphogypsum ([Fig. 5](#)).

Regarding the base saturation, the values were below the target of 70% in treatments with limestone in the 0–0.20 m layer, indicating that all limestone applied had reacted. The decrease in base saturation from 48 to 60 months is also an evidence of soil acidification. The highest values of base saturation as a function of phosphogypsum application without limestone were due to the increased calcium levels that were also observed in these treatments ([Fig. 5](#)). Similar effects were observed by [Soratto and Cruciol \(2008b\)](#).

#### 4.2. Soybean root growth

The soybean root characteristics were influenced only by the year effect ([Table 3](#)). Although the descriptions of the soybean cultivars are similar, the cultivars had variable reactions to the chemical properties of acid soils. This result explains the differences in root development between the two seasons because there was no water restriction during the development of culture. Despite the changes in the soil chemical characteristics, mainly in aluminium and exchangeable calcium, in the surface and subsurface layers, the application of lime and/or phosphogypsum ([Figs. 2 and 5](#)) did not affect the root growth or diameter. [Caires et al. \(2001\)](#) found similar results when evaluating soybean root growth with surficial liming with or without phosphogypsum.

Ca plays an important role in root growth, but the Ca critical limits for root growth appear to be very low. [Ritchey et al. \(1982\)](#) showed that even in soils with low Ca levels,  $1.0\text{--}1.5 \text{ mmol}_c \text{ dm}^{-3}$ , the seedling roots developed normally. In soybeans, root development occurred without any restrictions on soils with Ca of  $8.5 \text{ mmol}_c \text{ dm}^{-3}$  ([Rosolem and Marcello, 1998](#)). This effect may be one explanation for the observed absence of the root system growth response for soybean in the present study.

Toxic concentrations of Al cause a reduction in the growth rate of roots, but plant genotypes vary with respect to the reaction to Al present in acid soils. The Al toxicity of no-till systems is lower than that of conventional tillage ([Salet et al., 1996](#)). This lower toxic effect of Al in no-till systems has been attributed to the organic complexation of Al by soluble compounds present in the plant residues ([Franchini et al., 1999](#)).

#### 4.3. Plant nutrition

The nutrient concentrations ([Tables 4–6](#)) were within the ranges considered adequate for crops (soybeans, oats and sorghum) with the exception of the K and Zn levels for soybeans, the Ca, Mn and Fe levels for oat, and the Fe and Ca levels for sorghum, which were higher than the upper limit of these ranges ([Ambrosano et al., 1997](#); [Cantarella et al., 1997](#)).

The highest N, K, Mg, S, Zn and Fe concentrations in soybean were observed in 2008/2009 ([Table 4](#)). This result can be explained by the higher nutrient levels in the soil before planting (48 months after reapplication).

The increased N content in soybean and sorghum ([Tables 4 and 6](#)) demonstrates the beneficial effects of lime and/or phosphogypsum on the symbiotic  $\text{N}_2$  fixation process in soybeans, the nitrate availability in the soil and N mineralization

(Rosolem et al., 1990). The N<sub>2</sub> symbiotic fixation efficiency is associated with greater Mo and smaller Mn availability by reducing the soil acidity (Caires et al., 2000). Thus, the amount of N extracted by the soybean crop was positively correlated with the increasing Ca/Mn level in the leaves ( $r=0.91$ ,  $P<0.01$ ), corroborating the results of Caires et al. (2001). The availability of nitrate may increase with increasing pH through liming because the activity of nitrification can increase with increased pH in acidic soils (Islam et al., 2006).

The increases in Ca and Mg contents in crops due to liming in established no-till systems (Tables 4–6) have also been observed by several authors (Caires et al., 2006c, 2001; Castro and Crusciol, 2013). These effects have been attributed to the increases in the exchangeable Ca and Mg availability in the soil profile. Phosphogypsum, being a Ca source, can also increase the Ca uptake by crops (Caires et al., 2006b, 2001).

The greater absorption of S by soybean and sorghum (Tables 4 and 6) must have been caused by the release of the adsorbed sulphate or by the increased mineralization of organic matter as a consequence of the increased soil pH in the surface layers by liming (Fig. 3). Phosphogypsum application also increased the absorption of S by crops because it is an excellent S source.

The reduction in K absorption by soybean (Table 4) was accompanied by an increase in Ca and Mg in the leaves, showing the antagonistic effect of Ca and Mg on K ( $r=-0.66$ ,  $P<0.01$ ). Similar results were observed by Caires et al. (2008a). The micronutrient contents in soybean (Cu, Zn and Mn), oats (Mn and Fe) and sorghum (Mn) were also reduced with liming, with greater reductions observed in the phosphogypsum application for soybean (Cu, Zn and Mn) and black oat (Fe) (Tables 4–6). The increased soil pH with the use of limestone reduces the availability of some nutrients in the soil, particularly micronutrients (Malavolta et al., 1997). Sávio et al. (2011) also reported a reduction in the Zn and Mn concentrations in soybean grown in a no-till system with increased soil pH resulting from superficial liming.

#### 4.4. Yield components and grain yield

The soybean yield was higher in 2009/2010 due to the higher number of seeds per pod (Table 4), which is characteristic of the selected cultivar (CD 216). The shoot dry matter and yield components increased with liming (Tables 4–6). Phosphogypsum application increased the yield components in sorghum and soybean when associated with liming. However, black oat experienced a reduction in the number of panicles per m<sup>2</sup> with liming associated with phosphogypsum. These results reflect those observed for plant nutrition (Table 5). For sorghum, low temperatures extended the vegetative stage, and flowering occurred in a period with low water availability, mid-May 2010, which may have increased the spikelet sterility (Fornasieri Filho and Fornasieri, 2009).

The grain yield exhibited a strong correlation ( $P<0.01$ ) with population ( $r=0.69$ ) and 100-grain weight ( $r=0.84$ ) in soybean, panicles per m<sup>2</sup> ( $r=0.99$ ) in black oat, and panicles per m<sup>2</sup> ( $r=0.97$ ) and grains per panicle ( $r=0.98$ ) in sorghum (Tables 4–6). Similar results were observed by Sávio et al. (2011) and Corrêa et al. (2008) for soybean and by Soratto and Crusciol (2008c) for black oat.

In subtropical regions, grain production is not always benefited by liming, and the greatest yield improvements in the present study were obtained under water stress conditions. Thus, the importance of studies under different climatic conditions is evident, as in the present study in a tropical region under "cerrado" (tropical savannah), the main grain-producing region in Brazil, liming consistently benefited grain production, even in years with regular rainfall distribution.

## 5. Conclusion

Superficial liming with or without phosphogypsum reduced the surface and subsurface soil acidity 5 years after application in a no-till system. The movement of Ca<sup>2+</sup> and Mg<sup>2+</sup> from the surface layer into the subsoil over time was evident. Phosphogypsum application associated with liming increased the Ca<sup>2+</sup> levels throughout the soil profile. Liming maintained high levels of Mg<sup>2+</sup> throughout the soil profile with or without phosphogypsum application. The organic matter content increased with liming with or without phosphogypsum, indicating that in the long term, these practices can increase the C accumulation in the system. Phosphogypsum application had a residual effect on the SO<sub>4</sub>-S levels, and high sulphate concentrations were observed in the subsoil after 5 years. Superficial liming improved crop nutrition and, when associated with phosphogypsum, increased Ca absorption by soybean and sorghum, as reflected in the increased productivity of these crops.

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